

# Principles of Radiative Transfer; Biosphere and Land surface Processes

## Joint IMD-WMO Group Fellowship Training on Numerical Weather Prediction

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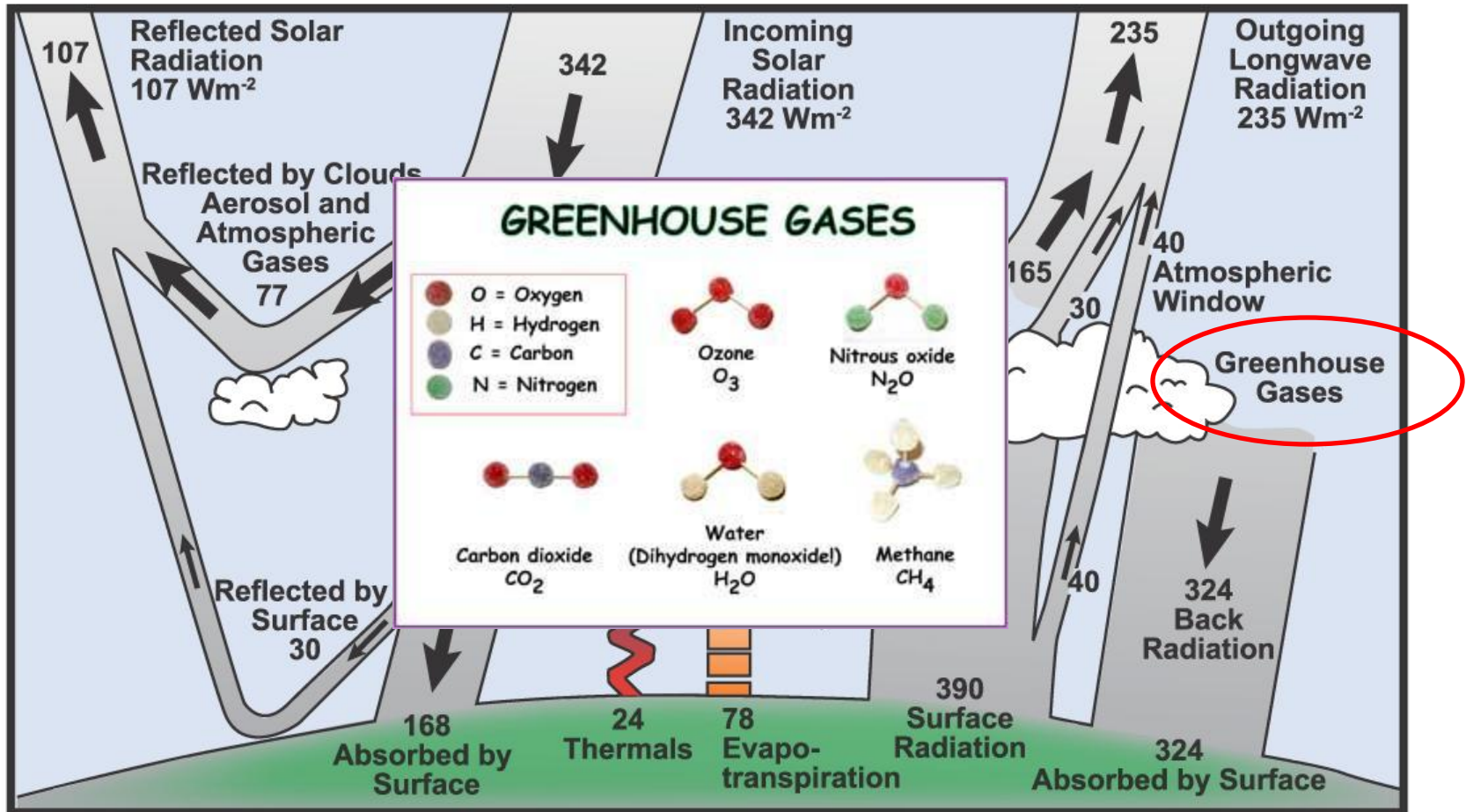
&

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Sciences  
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India*



# Earth's energy

Source: IPCC

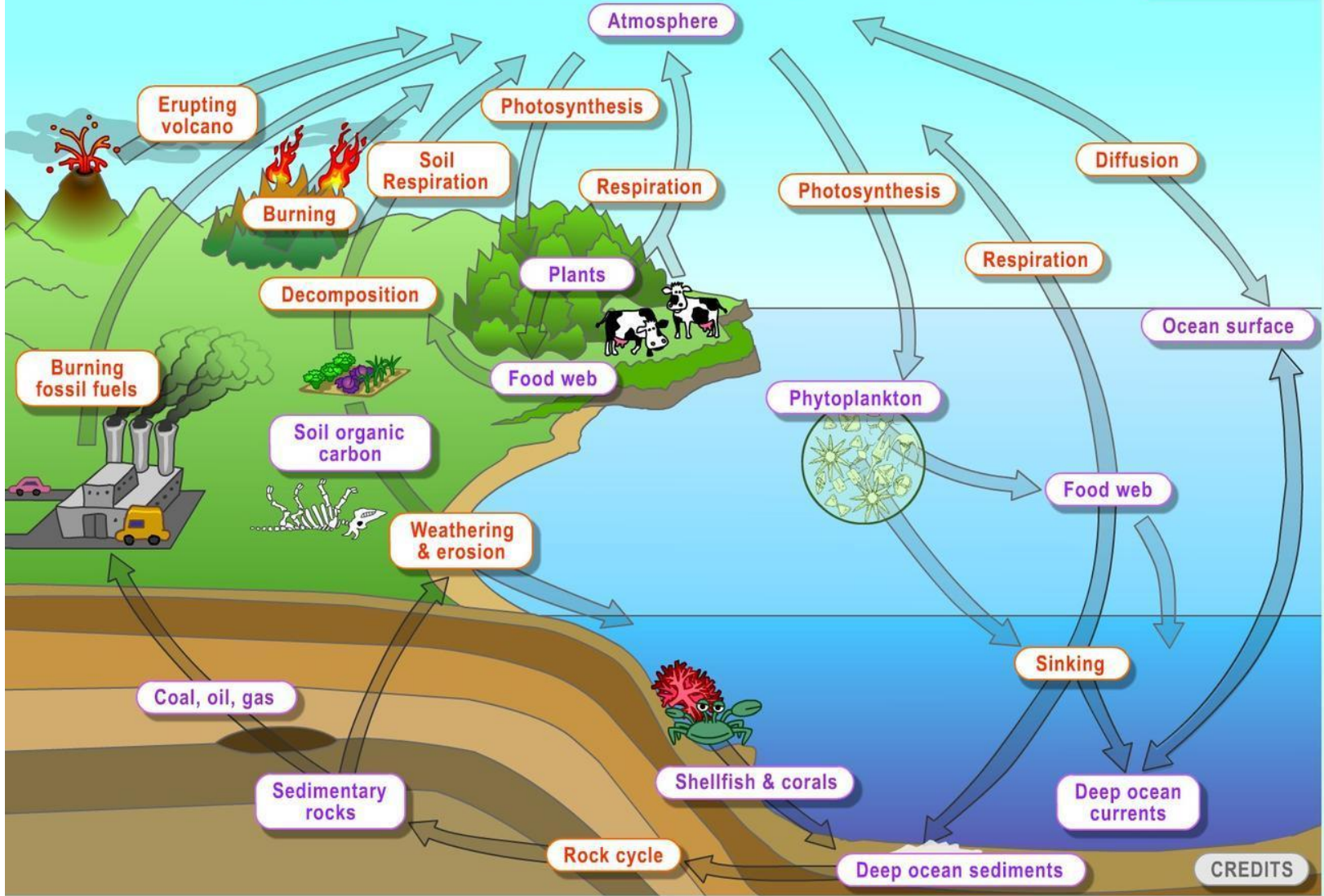


at surface the energy budget is,

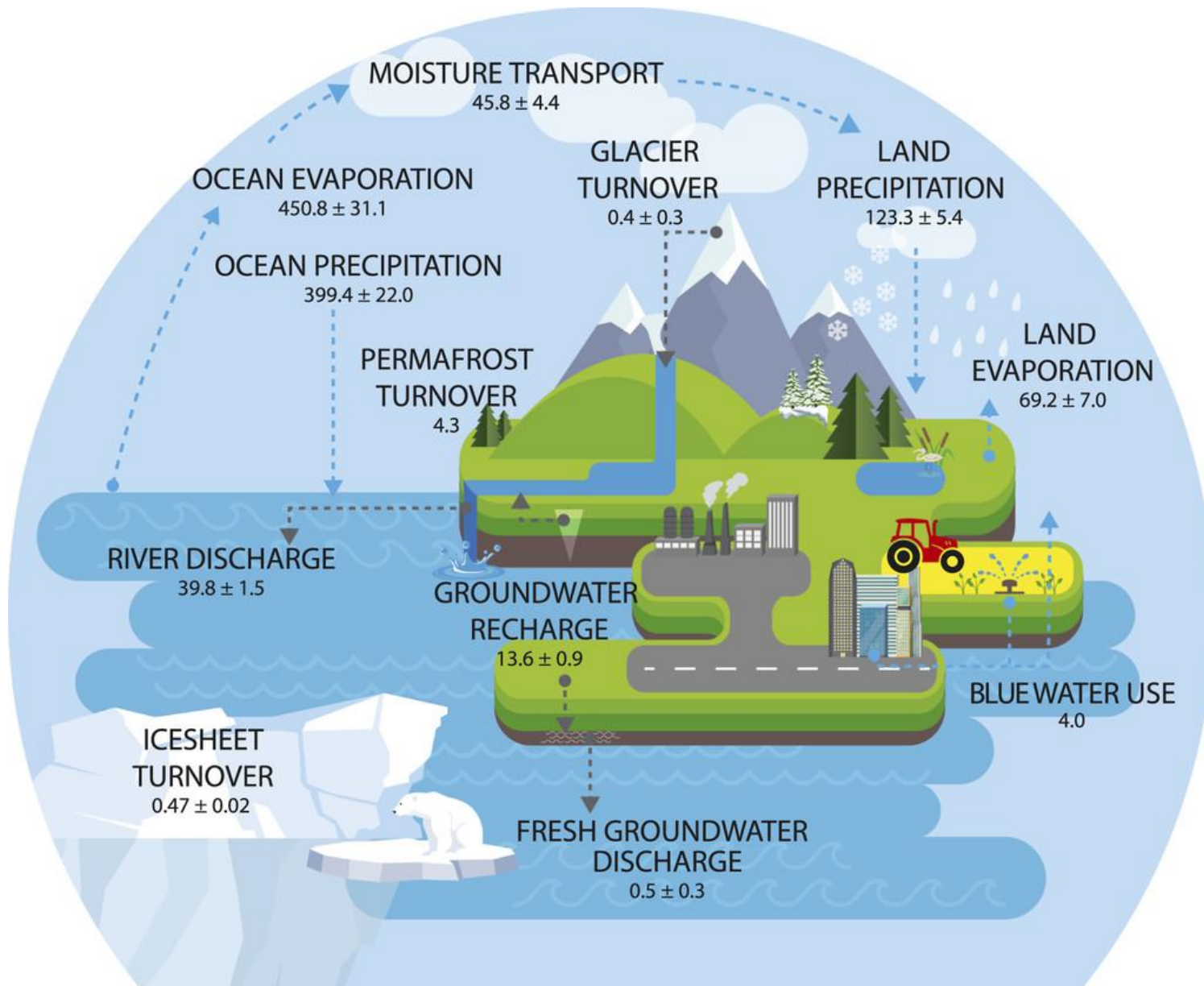
$$R_n = R_{SW}(in) + R_{LW}(in) - R_{SW}(out) - R_{LW}(out), \quad \longrightarrow \quad R_n = H + LE + G.$$

# THE CARBON CYCLE

KEY	Process
	Reservoir



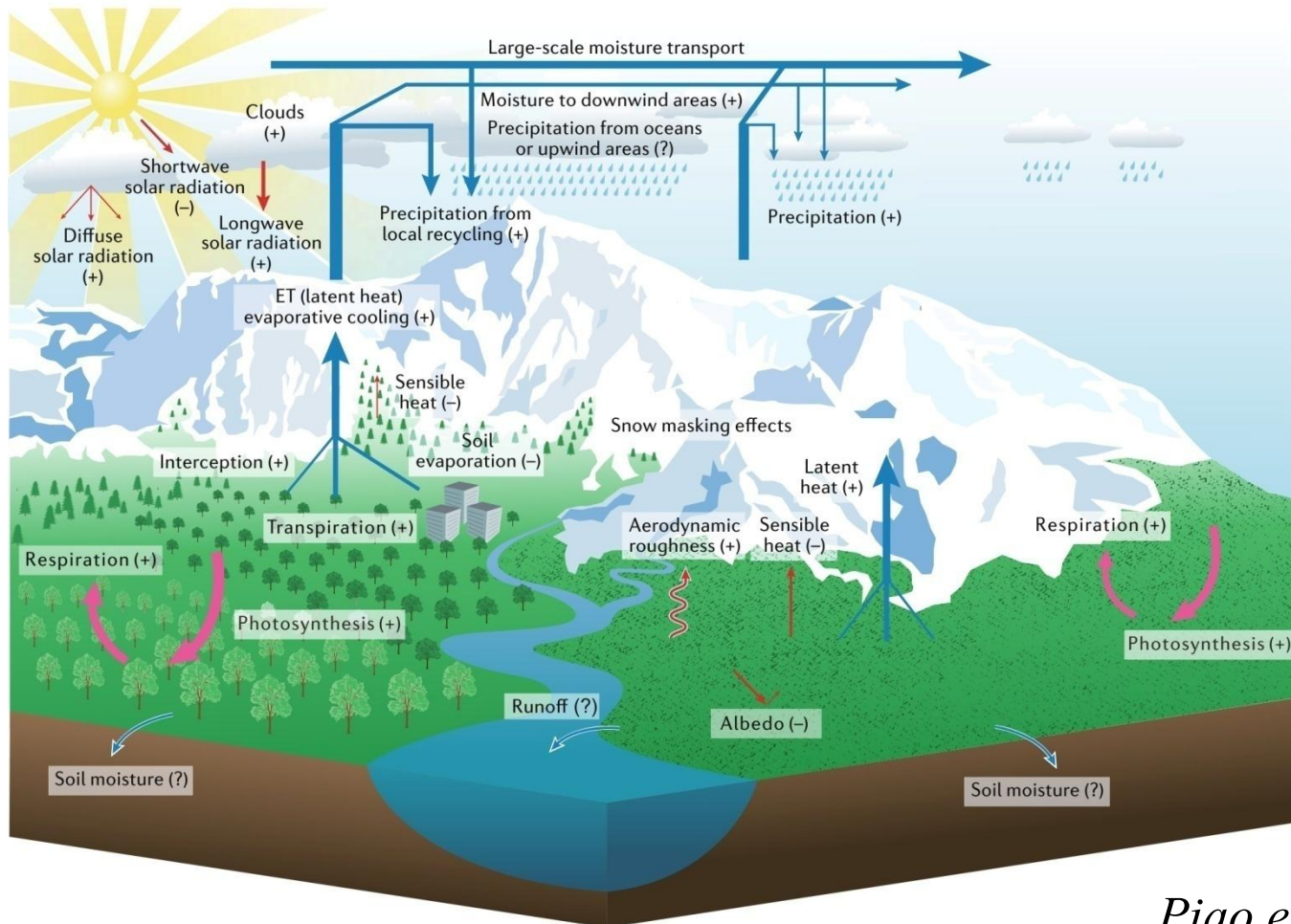
Picture courtesy: [www.cleanet.org](http://www.cleanet.org)



## GLOBAL WATER CYCLE FLUXES

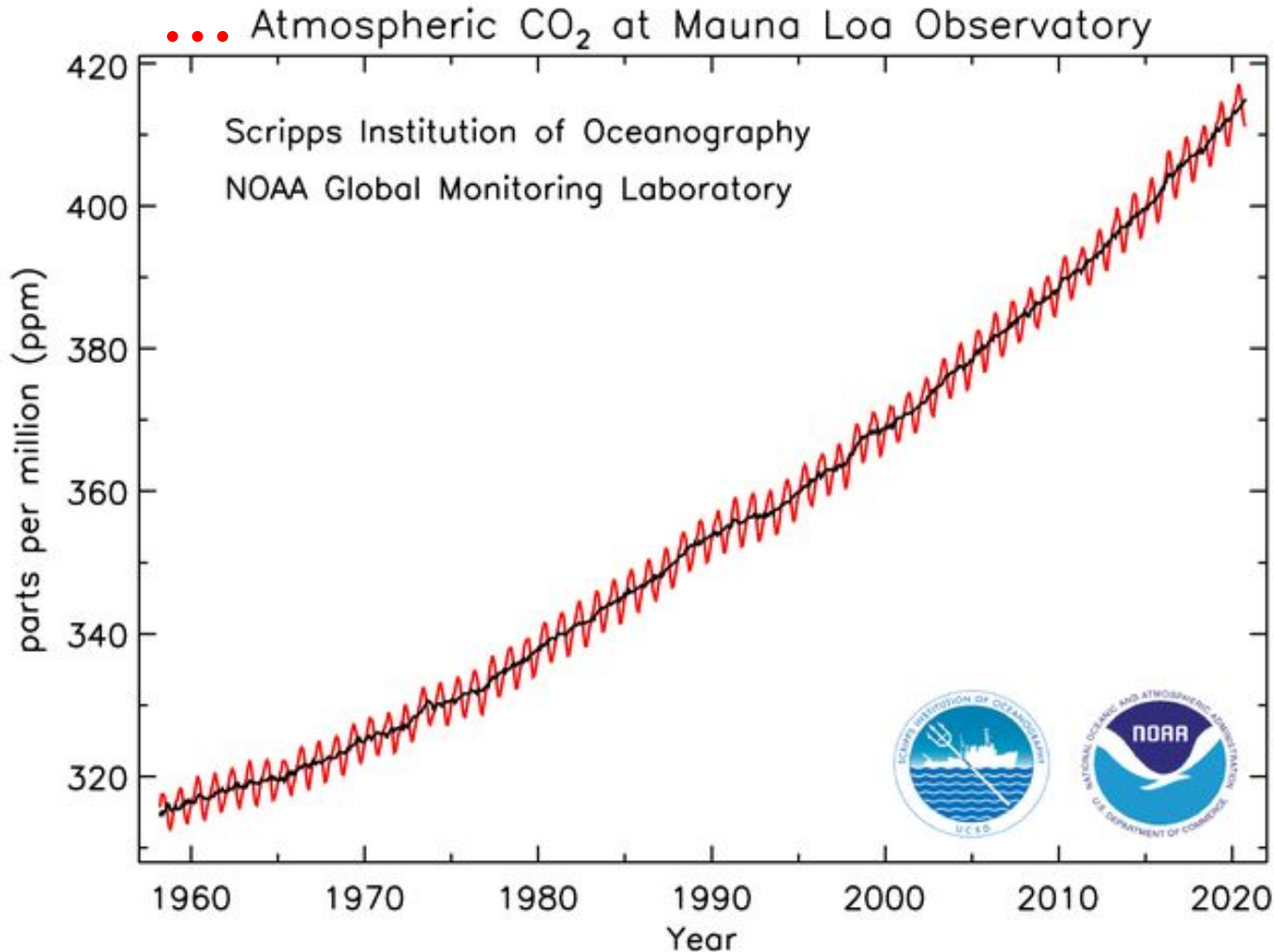
# vegetation-atmosphere feedback

Schematic diagram showing how vegetation greening affects climate via land-atmosphere exchange of carbon, water and energy. For each process or flux, the symbols '-', '+' and '?' in brackets represent a decreasing, increasing and unknown trend, respectively, in response to vegetation greening. Vegetation greening cools global climate mainly through enhancing land carbon sink and evapotranspiration.



*Piao et al., 2020*

# Increasing CO<sub>2</sub> in the atmosphere

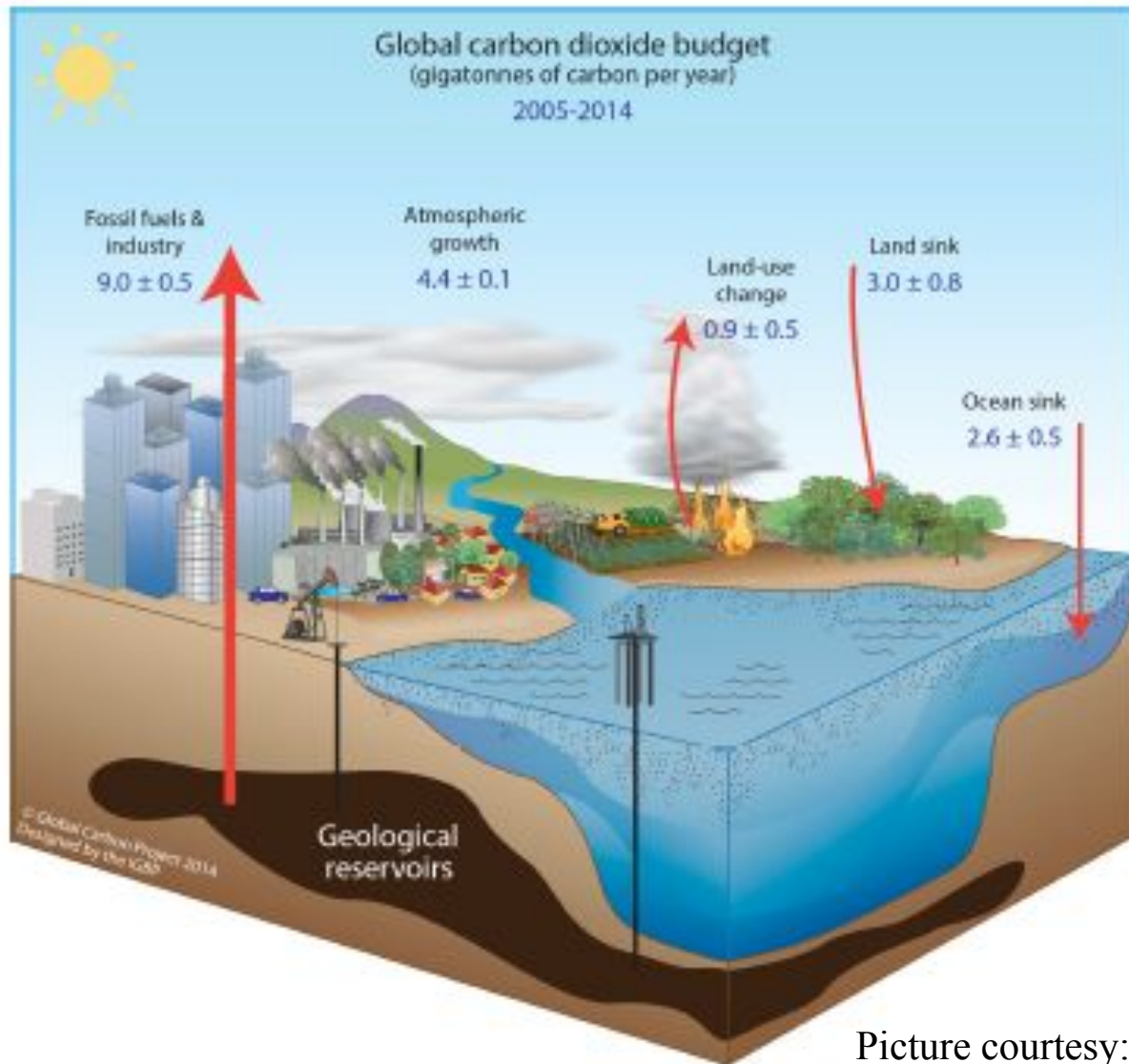


**present day  
~ 414 ppm**

**pre-industrial  
~ 280 ppm**

Source: [Global Monitoring Laboratory - Carbon Cycle Greenhouse Gases \(noaa.gov\)](https://www.noaa.gov/global-monitoring-laboratory-carbon-cycle-greenhouse-gases)

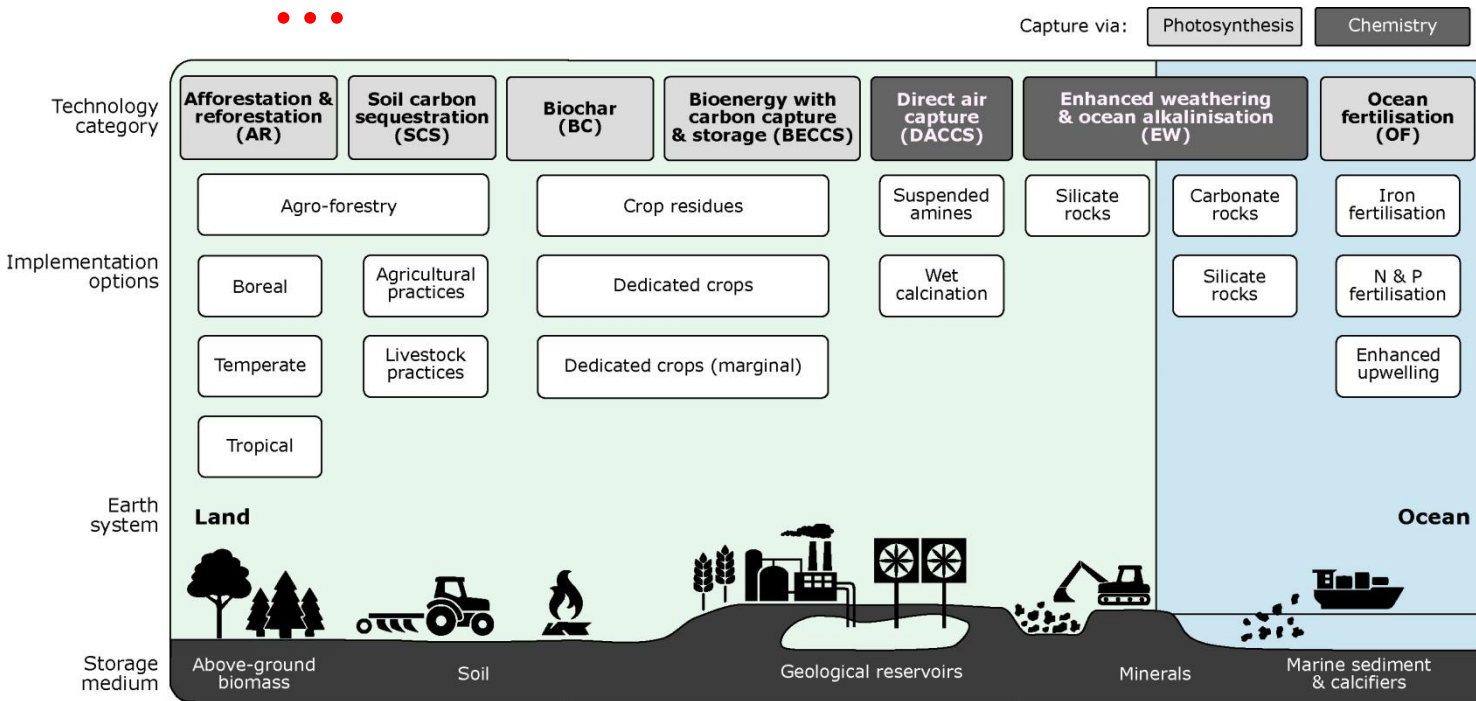
# Why is it important to study



- ❖ Terrestrial Ecosystems are a major sink of atmospheric CO<sub>2</sub>
- ❖ During 2009-2018,  
 $S_{\text{land}} = 3.2 \pm 0.6 \text{ GtC yr}^{-1} > S_{\text{ocean}} = 2.5 \pm 0.6 \text{ GtC yr}^{-1}$  (Friedlingstein et al., (2019))
- ❖ Paris climate accord: limit global warming to 2 °C in the present century (Source: UNFCCC)
- ❖ Proper accounting of carbon budgets are required to estimate the *intended nationally determined contribution* (INDC) and devise mitigation strategies

Picture courtesy: Le Quéré et al., (2015)

# In the context of climate change



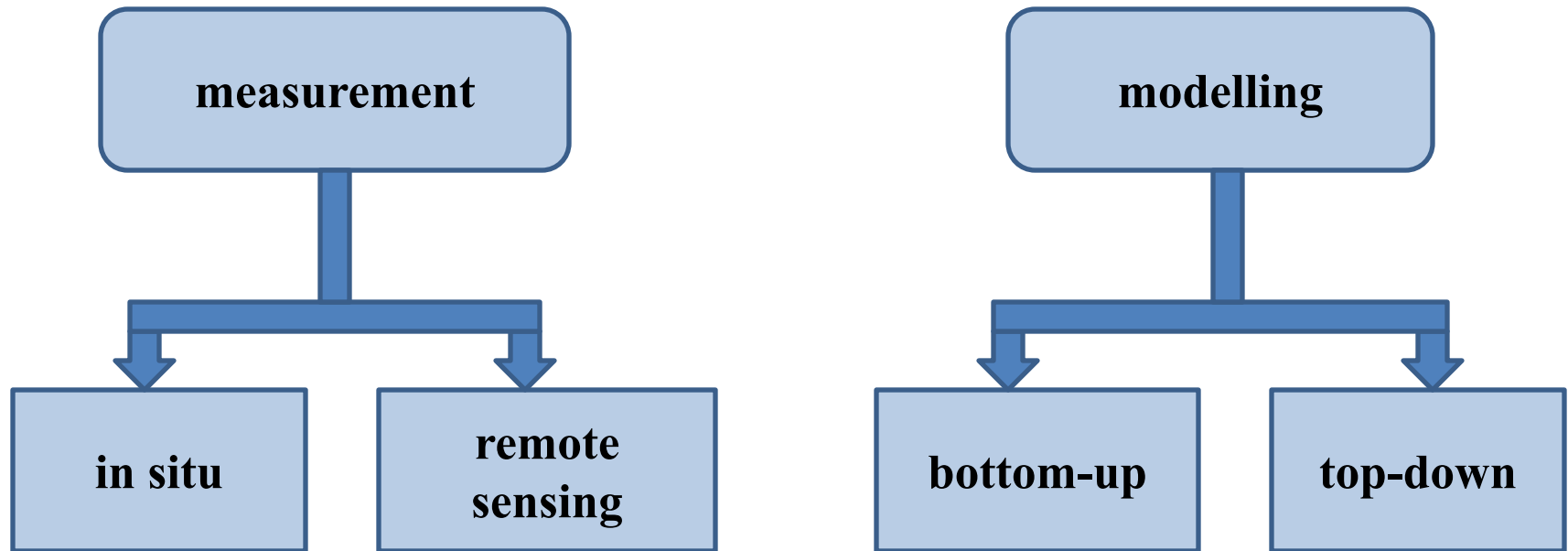
Several mitigation strategies are based on land

Picture courtesy : Minx et al., (2018)

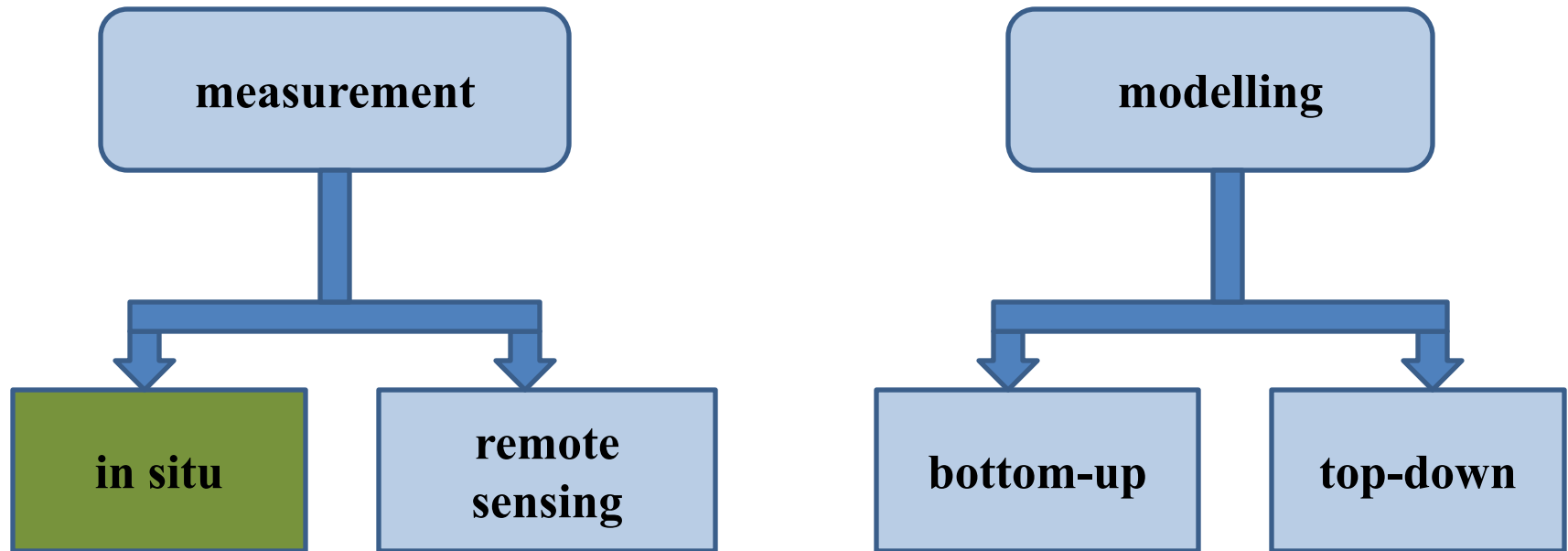
- ❖ However, large-scale implementation can result in loss of biodiversity in terrestrial ecosystems, including coastal and wetland (IPBES report; Bongaarts, (2019)
- ❖ Climate change puts additional burden on terrestrial ecosystems, including food and fibre production to cater a growing population etc. in all SSPs (Shared Socio-economic Pathway) (IPCC Special Report on Climate Change and Land).
- ❖ Climate Change or Food Security?
- ❖ Fresh potable water resources are also heavily burdened.



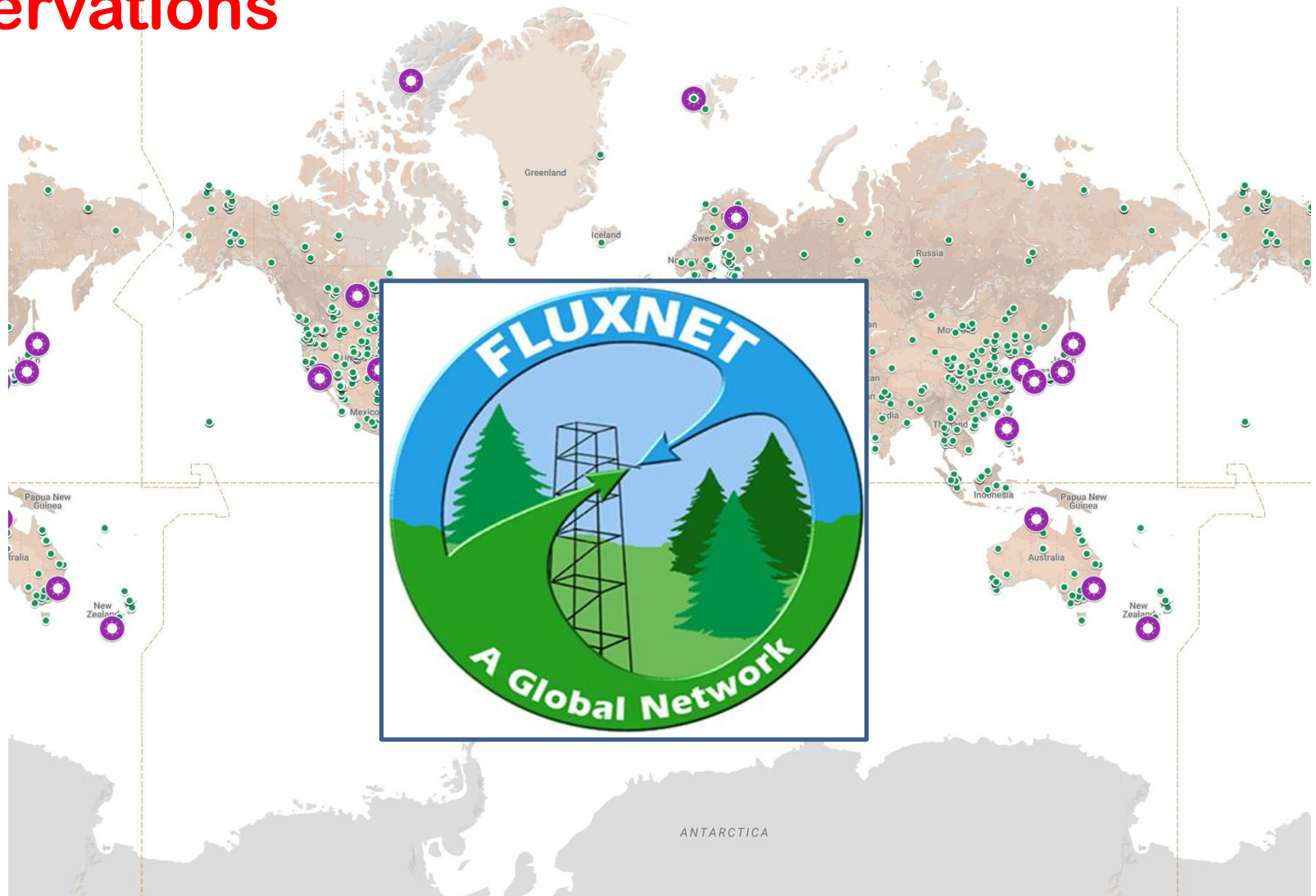
# How to estimate the biosphere-atmosphere flux exchanges?



# How to estimate the biosphere-atmosphere flux exchanges?



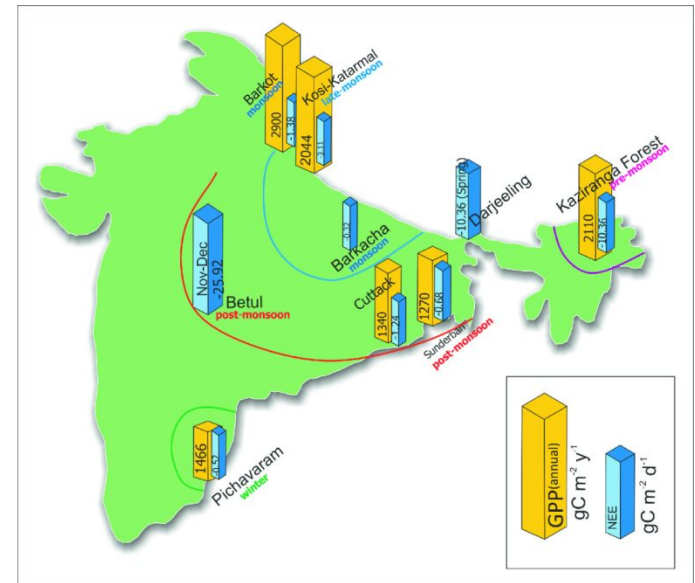
# Global ecosystem-atmosphere flux observations



# In the Indian context ...



1. Wide diversity of plant, forest types and climatic regimes
2. Poorly represented in terrestrial ecosystem, Earth system, and climate models
3. Integrated data-model framework required



Source: India State of Forest Report 2019, Ministry of Environment, Forest & Climate Change, Government of India.

Picture courtesy: Assessment of Climate Change over the Indian Region 2020, Ministry of Earth Sciences, Government of India.



Canopy Crane, Wind River, Washington, USA

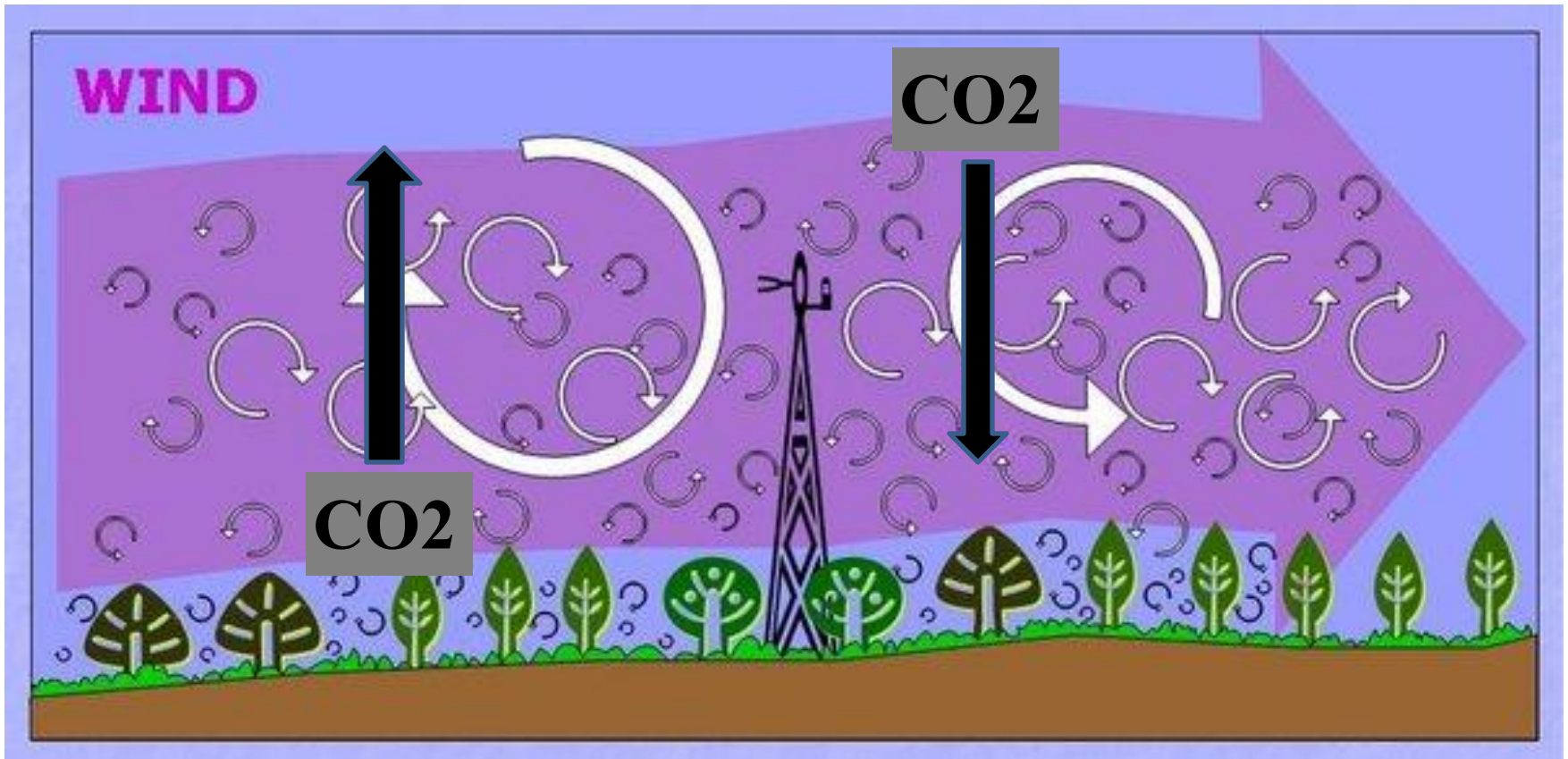
Picture Courtesy: Wikipedia

wind parameters  
gas (CO<sub>2</sub>, water vapour, CH<sub>4</sub> etc.) concentrations  
weather parameters (air temperature, pressure, humidity, precipitation etc.)  
radiation (incoming and outgoing shortwave and longwave)  
photosynthetically active radiation (visible, 40-700 nm)  
soil parameters (temperature, moisture, heat flux etc.)  
plant physiology, leaf area index (LAI) etc.



MetFlux India flux tower at the Pichavaram mangroves, Tamil Nadu, India

Picture courtesy: Deb Burman, (2020b)



Eddies are random turbulent motions of wind; these result in the spatial movement of air parcels and in the process, the gas molecules.

Picture courtesy: Dr. George Burba, Li-COR Biosciences, USA

# Let's do some mathematics ....

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Section

Abstract

Experiment

**XXIX. An experimental investigation of the circumstances which determine whether the motion of water shall be direct or sinuous, and of the law of resistance in parallel channels**

Osborne Reynolds

Published: 01 January 1883 | <https://doi.org/10.1098/rstl.1883.0029>

### Reynolds averaging:

$$x = \bar{x} + x'$$

$\bar{x}$  = time mean; usually taken for half an hour

$x'$  = instantaneous fluctuations around the mean

- Need to sample the eddies of all size
- Cumulative contribution to the net exchange
- Flux = mass (or energy) transferred per unit area per unit time
- more eddies sampled, better accurate is the measurement. ho, high-frequency measurements are required, usually 5, 10 or 20 Hz

# In vertical dimension ....

- in a turbulent flow (such as in atmosphere),
- the vertical flux of any species  $c$ ,  $F_c = \rho_a \overline{ws}$
- where  $s = \rho_c / \rho_a$
- $\rho_c$  = density of species  $c$  ( $\text{CO}_2$ , in our case)
- $\rho_a$  = density of dry air
- using Reynolds averaging,

$$\rho_c = \overline{\rho_c} + \rho_c'$$

$$w = \overline{w} + w'$$

$$\rho_a = \overline{\rho_a} + \rho_a'$$

- Replacing and simplifying,



- $F_c = \overline{(\bar{\rho}_c + \rho_c')(\bar{w} + w')(\bar{\rho}_a + \rho_a')}$
- expand the terms in parentheses
- simplify the expression...
- $F_c = \bar{\rho}_a \overline{w's'}$

- $$F_c = \overline{(\bar{\rho}_c + \rho_c')(\bar{w} + w')(\bar{\rho}_a + \rho_a')}$$

- expand the terms in parentheses
- simplify the expression...

- $F_c = \bar{\rho}_a \bar{w} \bar{s}$  mind the gap (and the primes) !

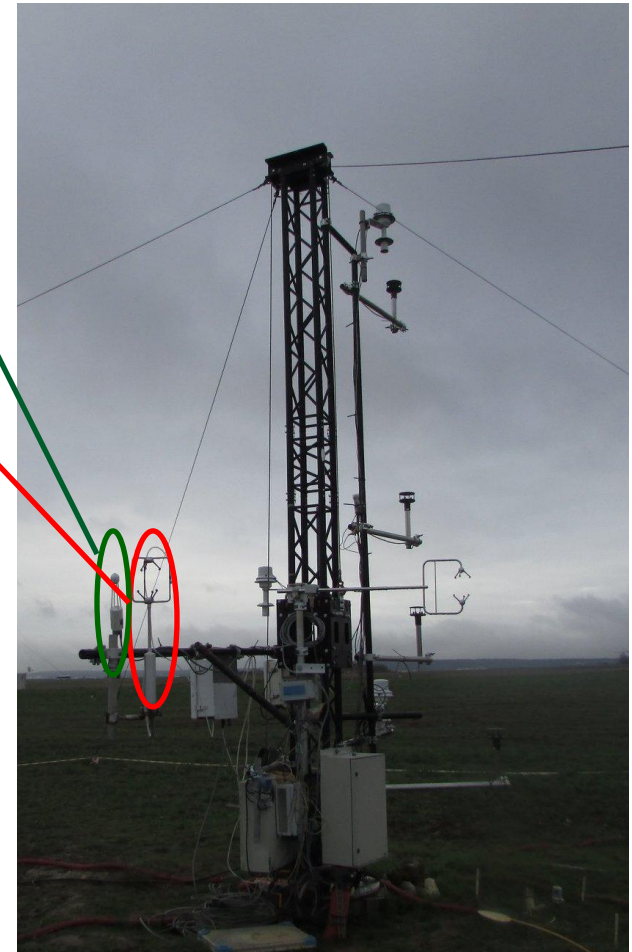
- actually, not so simply (Aubinet et al., (2012))

# Quality Control

1. despiking to remove erroneous spikes (Vickers & Mahrt, 1997)
2. detrending to remove any long-term trend or ramp in the measurement (Kaimal & Finnigan, 1994)
3. coordinate rotation to ensure proper measurement geometry (Kaimal & Finnigan, 1994)
4. WPL or density correction to account for the atmospheric moisture (Webb et al., 1980)
5. time-lag corrections to ensure the measurements of wind and gas concentrations are synchronous (Moncrieff et al., 1997, 2004)
6. and many more ... read Aubinet et al., (2012)

gas

wind



Eddy Covariance flux tower, France

- Why do we need moisture correction?



- Why do we need moisture correction?

moist air is lighter than dry air. If not corrected for the moisture, the measured gas densities will be higher than actual.

# We learnt about the Eddy Covariance method; There are other methods as well

1. Flux-Gradient (Lee, 2018)  
mass flows from high  
density to low density

$K_c$  is the eddy diffusivity factor  
for  $\text{CO}_2$

$$F_c = -K_c \frac{\partial \bar{c}}{\partial z}$$
$$= -K_c \frac{\overline{c(2)} - \overline{c(1)}}{z(2) - z(1)}$$

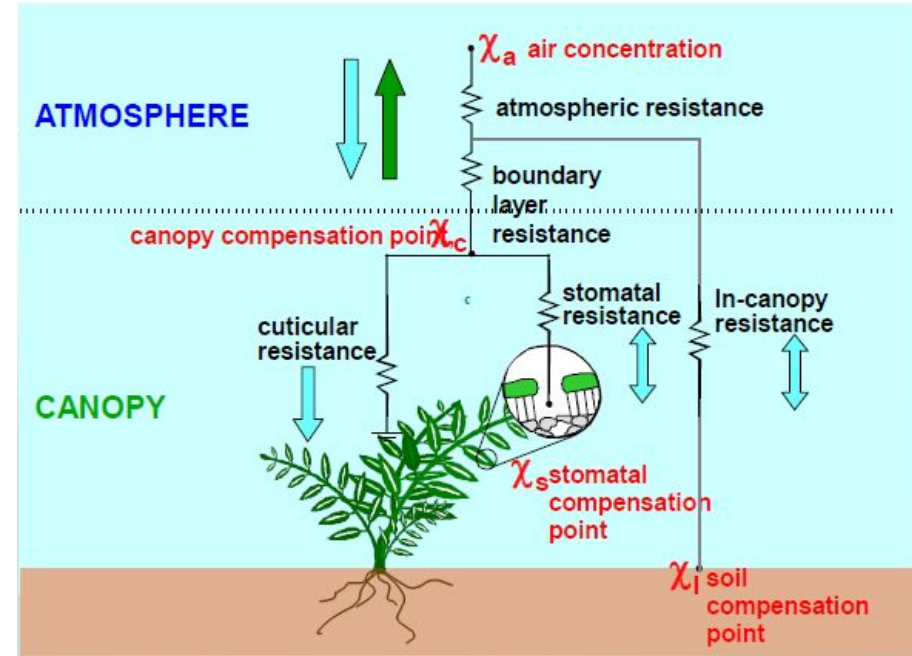
# We learnt about the Eddy Covariance method;

There are other methods as well

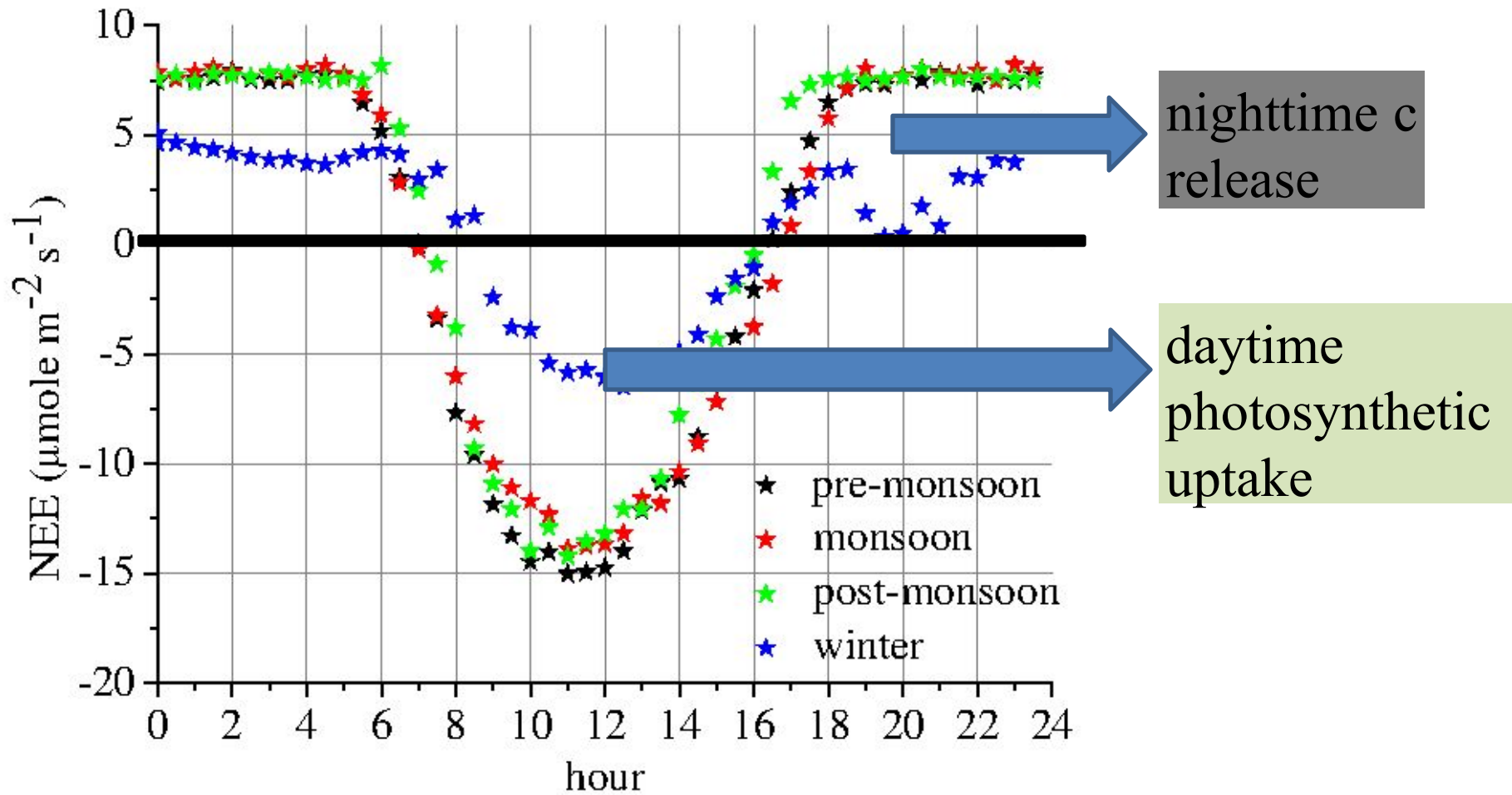
## 2. Resistance formulation (analogous to the concept of flow of electricity or fluid)

Nemitz et al. (2000)

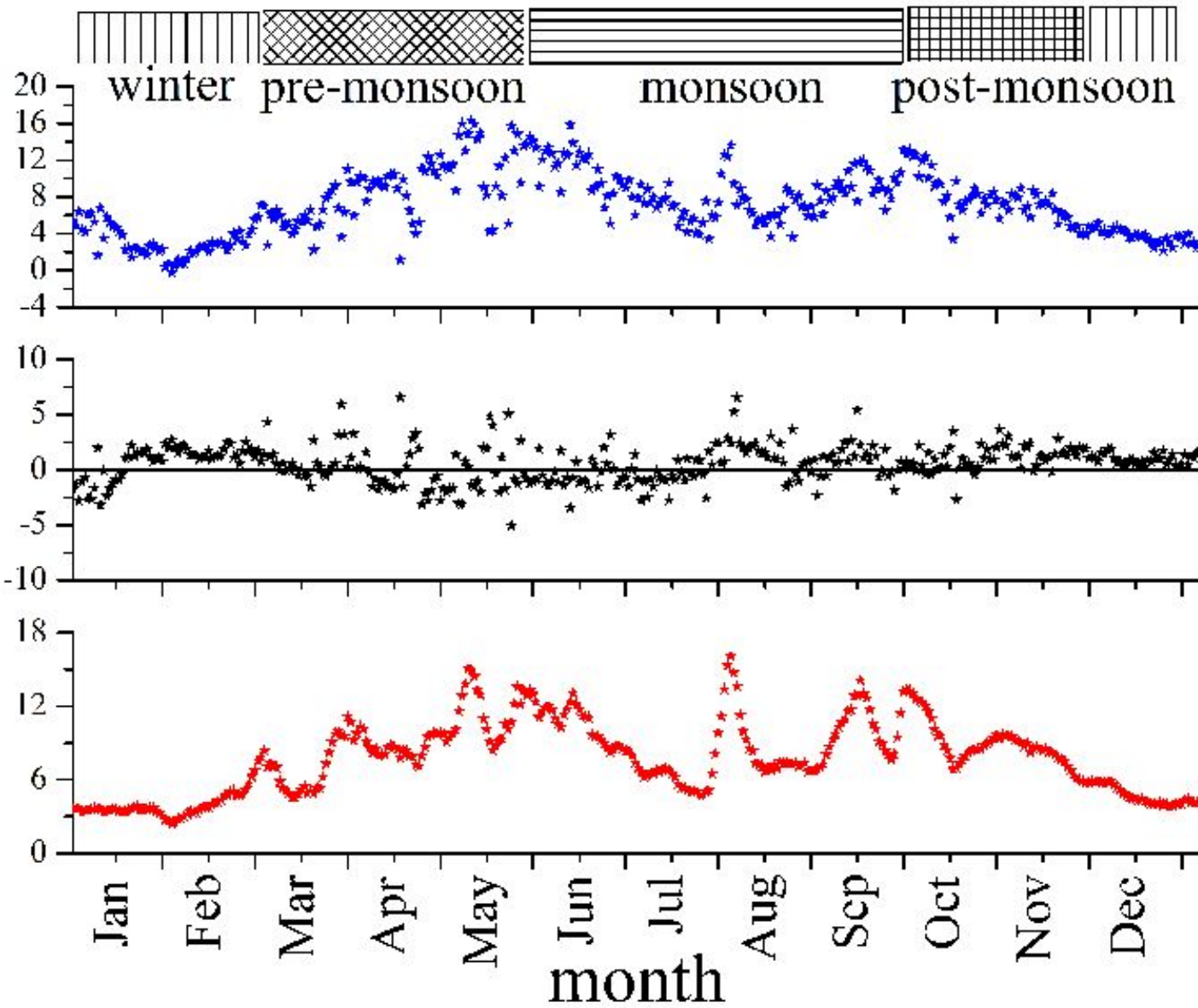
Bigleaf assumption: entire canopy is  
considered to behave as a big leaf  
(Sellers et al., 1992)



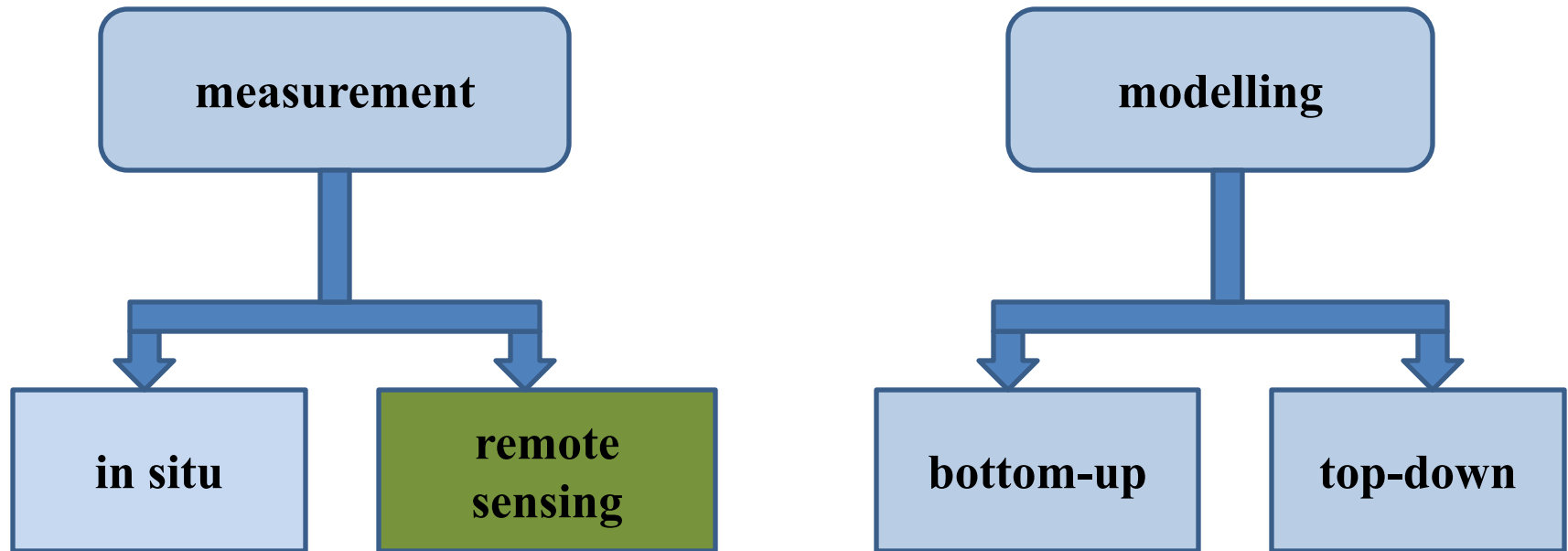
$$\text{Stomatal Resistance } r_s = \frac{1}{m \frac{A_n \cdot h_s}{c_s} + b}$$







# How to estimate the biosphere-atmosphere flux exchanges?

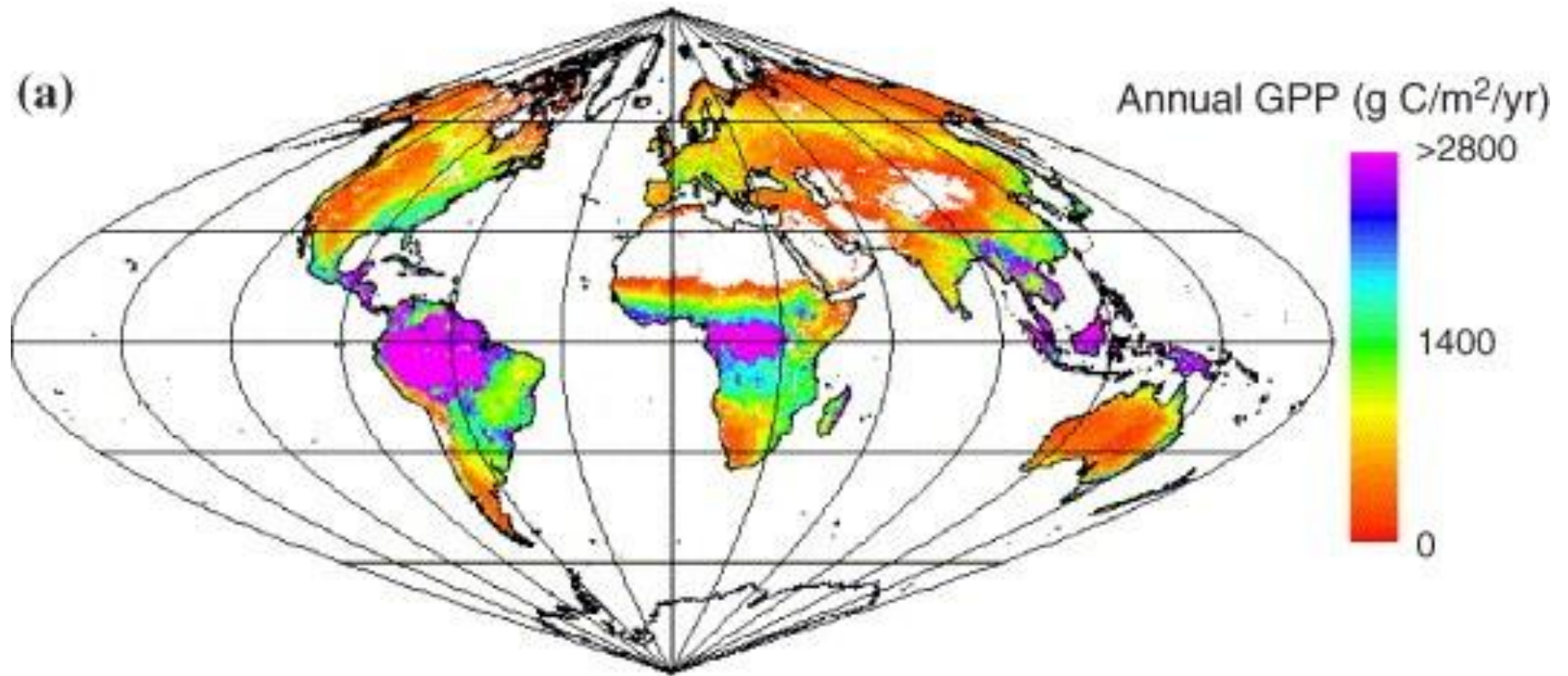


## Vegetation Indices (VI):

1. ratio of reflectances in different band (Huete et al., 2002)
2. Normalised difference vegetation index (NDVI),

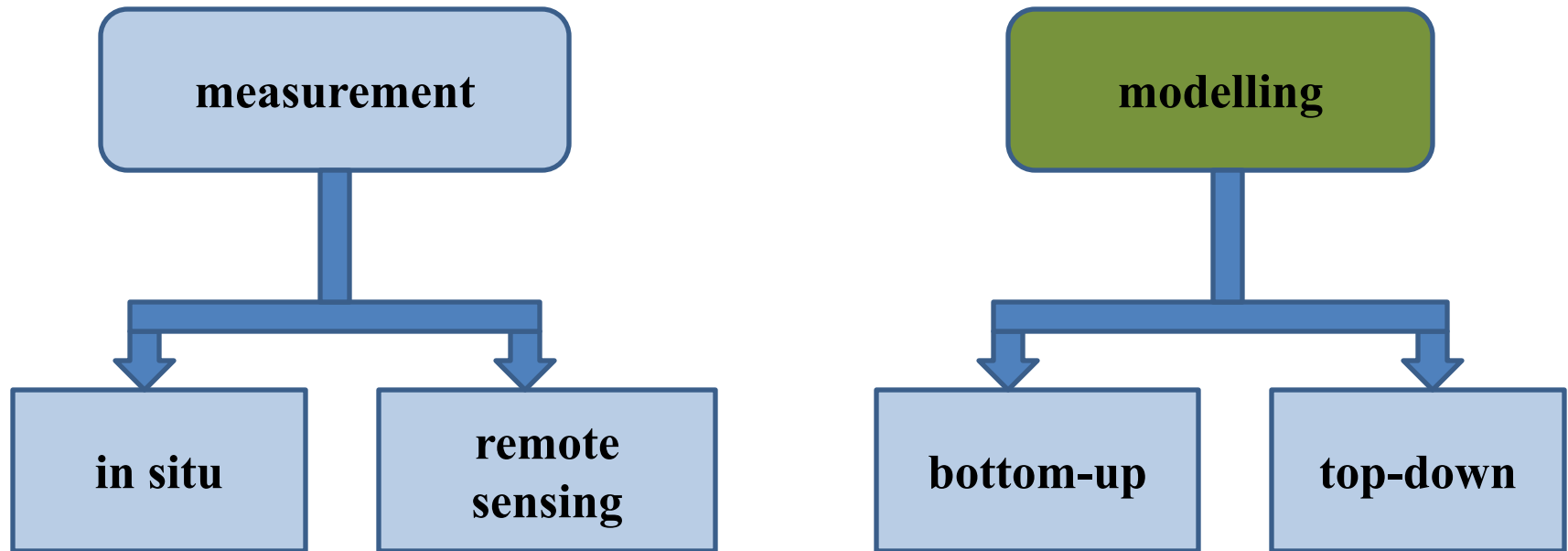
$$NDVI = \frac{\alpha_{nir} - \alpha_{vis}}{\alpha_{nir} + \alpha_{vis}}$$

3. Leaf-Area Index (LAI): total one-sided leaf area per unit ground area (Watson, 1947)
4. Enhanced Vegetation Index (EVI), Land-surface Water Index (LSWI) etc.



mean global GPP(gC m<sup>-2</sup> y<sup>-1</sup>) distribution, estimated from MODIS (Zhao et al., (2005))

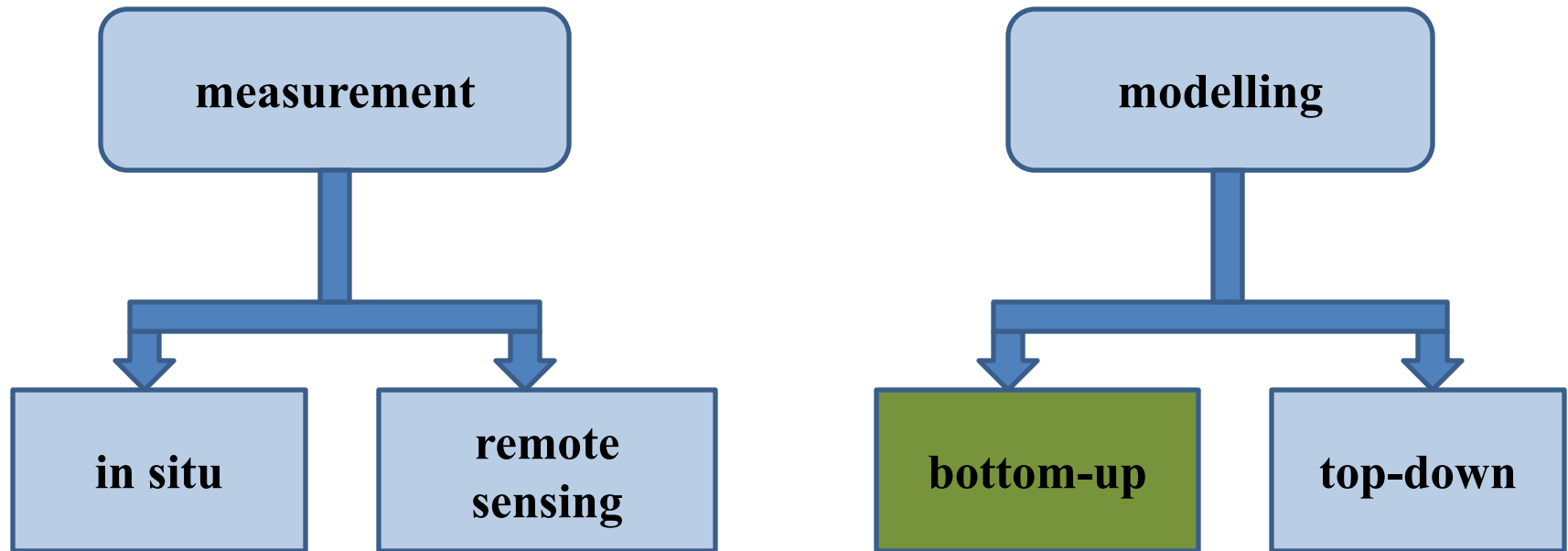
# How to estimate the biosphere-atmosphere flux exchanges?



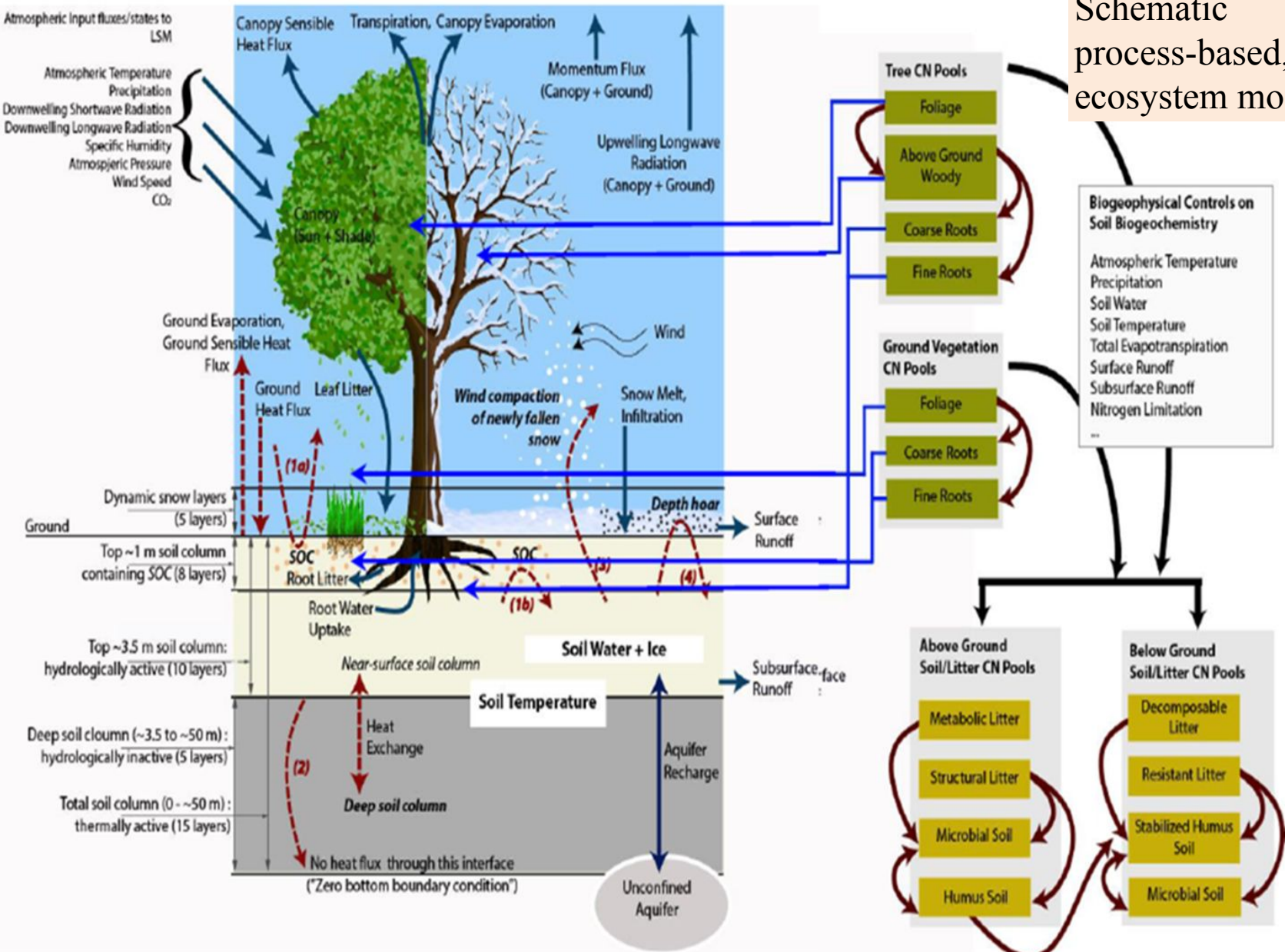
# What is a model?

1. a set of mathematical (differential, integral, algebraic etc.) equations
  - equations of motion
  - continuity equation
  - mass conservation
  - thermodynamic equations
2. solved at single (point model) or multiple grid points (regional and/or global)
3. the governing equations are derived from our knowledge of physics, chemistry, biology, hydrology etc.
4. if that does not work, we parameterize in phenomenological way, i.e. based on the knowledge gained from observations.
5. in short, we recreate the complex coupled interplay among meteorology, physiology, hydrology etc. to study the evolution of the system and predict its response

# How to estimate the biosphere-atmosphere flux exchanges?



# Schematic of a process-based, bottom-up ecosystem model



NOAH LSM (Land Surface Model), JULES (Joint UK Land-Environment Simulator), CESM (Community Earth System Model) etc.

Picture Courtesy: Prof. Atul K. Jain, Department of Atmospheric Science, University of Illinois at Urbana-Champaign USA



**Table A2**  
ISAM equations for the carbon modules.

Function	Equations
Plant autotrophic respiration	$R_{\text{leaf}} = k_{\text{leaf}} \times \frac{C_{\text{leaf}}}{\text{CN}_{\text{leaf}}} \times \varnothing \times gt$ (S1)
	$R_{\text{stem}} = k_{\text{stem}} \times \frac{C_{\text{stem}}}{\text{CN}_{\text{stem}}} \times gt$ (S2)
	$R_{\text{root}} = k_{\text{root}} \times \frac{C_{\text{root}}}{\text{CN}_{\text{root}}} \times \varnothing \times gt$ (S3)
	$Q_{10} = 3.22 - 0.046 \times T_{\text{veg}}$ for leaves and stems (S4)
	$Q_{10} = 3.22 - 0.046 \times T_{\text{soil}}$ for roots (S5)
	$gt = Q_{10}^{(T_{\text{veg}} - 20) / 10}$ for leaves and stems (S6)

### Photosynthetic pathways:

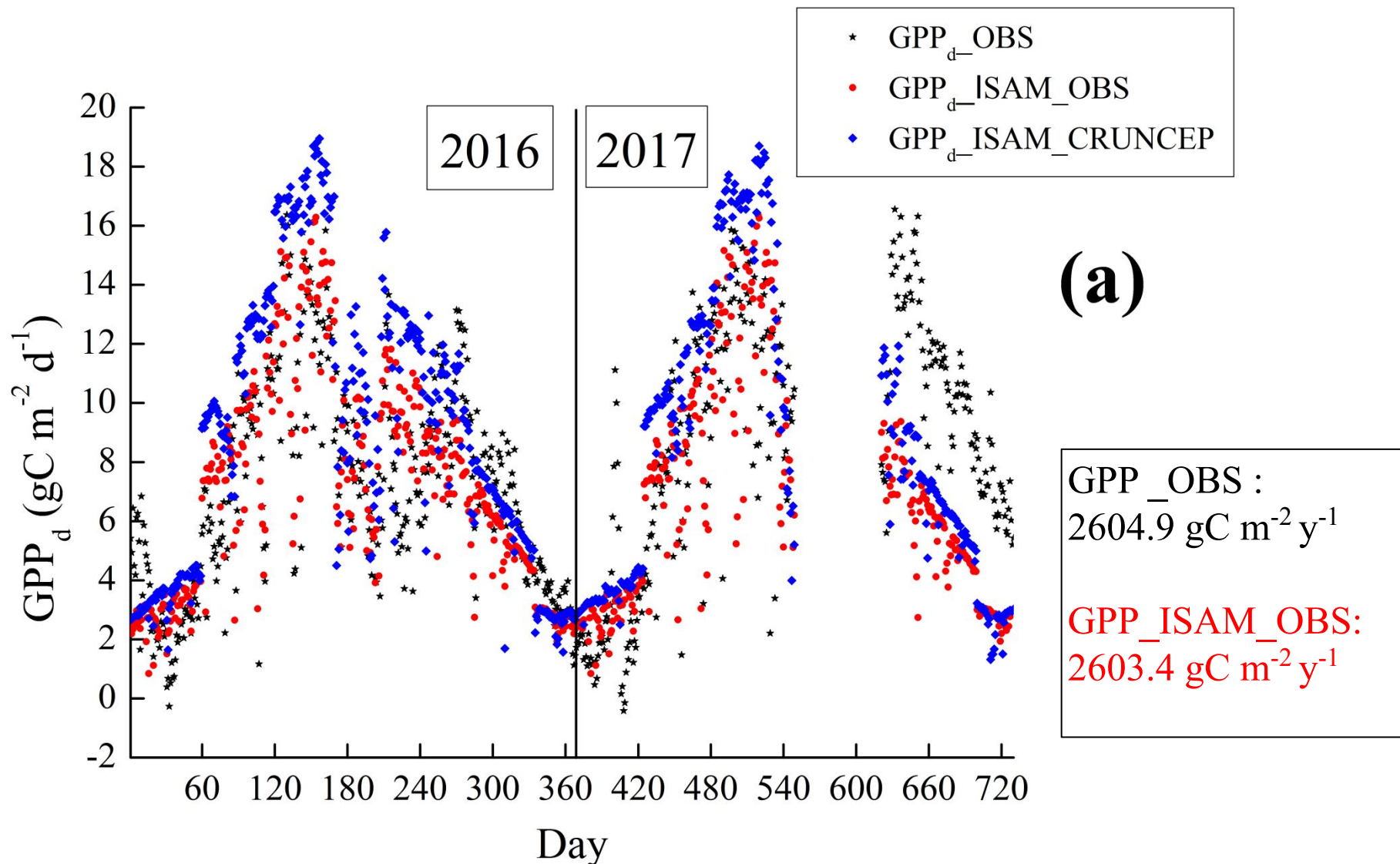
1. C3 (Farquhar et al., 1980; Collatz et al., 1991);  
Maximum carboxylation capacity ( $V_{\text{cmax}}$ ), Bulk canopy  
maximum electron transport rate ( $J_{\text{max}}$ )
2. C4 (Collatz et al., 1992)
3. CAM (Cortázar & Nobel, (1990); Kluge & Ting,  
(2012))

*El-Masri et al., 2013*

### Respiration:

1. autotrophic respiration (plants) + heterotrophic  
respiration (animals, microbes)
2. maintenance + growth; usually taken as a  
percentage of photosynthetic uptake

# calibrating and validating a process-based ecosystem model

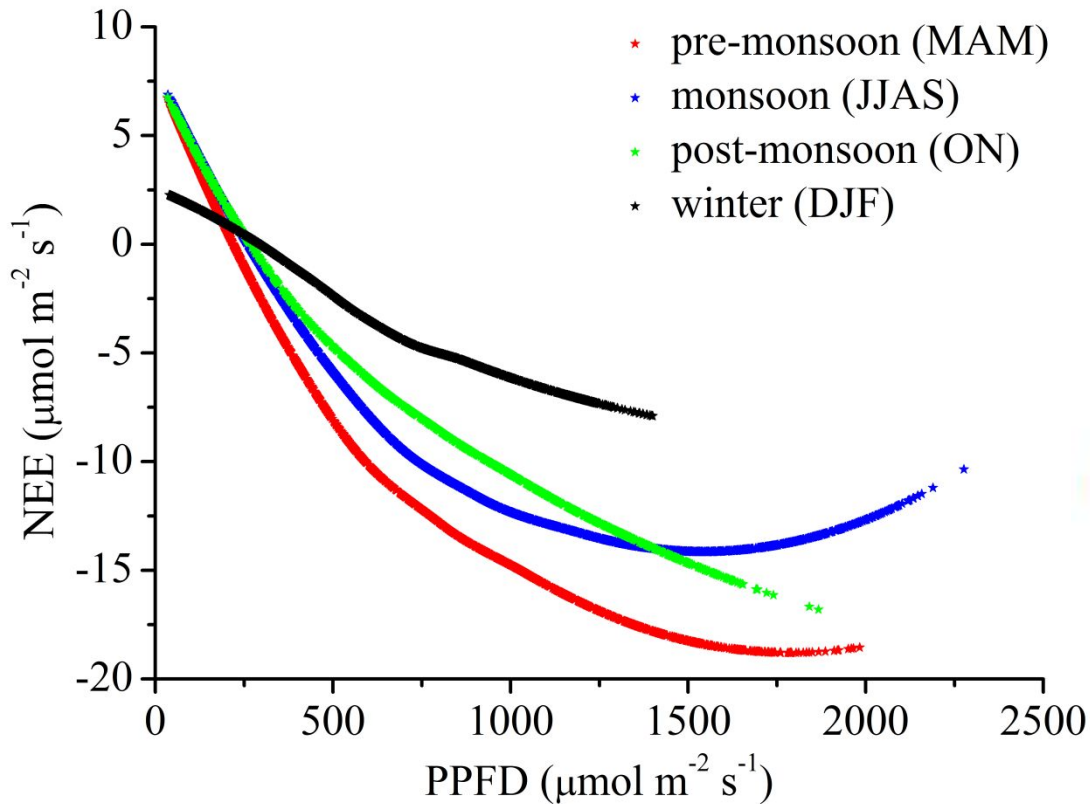


**functional relationships  
derived from observations are  
represented in models**

Michaelis-Menten relation:

$$NEE = \frac{\alpha \cdot PPFD \cdot NEE_{sat}}{\alpha \cdot PPFD + NEE_{sat}} + TER.$$

*Deb Burman et al., 2020*



Season	$\alpha$ ( $\mu\text{mol CO}_2$ $\mu\text{mol}^{-1}$ photons)	$NEE_{sat}$ ( $\mu\text{mol}$ $\text{m}^{-2} \text{s}^{-1}$ )	$F_m$ ( $\mu\text{mol m}^{-2}$ $\text{s}^{-1}$ )	LCP ( $\mu\text{mol}$ $\text{m}^{-2} \text{s}^{-1}$ )	$R_d$ ( $\mu\text{mol m}^{-2}$ $\text{s}^{-1}$ )
pre-monsoon	- 0.05	- 47.48	- 20	250	7.5
monsoon	- 0.05	- 35.33	- 15	250	10
post-monsoon	- 0.04	- 41.31	- 20	250	10
winter	- 0.02	- 41.23	- 10	250	2.5

# Light Use Efficiency (LUE) approach

- ❖ Proposed first by Monteith, 1972
- ❖ Carbon uptake by a canopy is directly proportional to the absorbed photosynthetically active (visible band, 400-700 nm) radiation (PAR)
- ❖ Proportionality constant ( $\epsilon$ ) is the light use efficiency. It depends on the canopy type, vegetation health, moisture and nutrient stress etc.

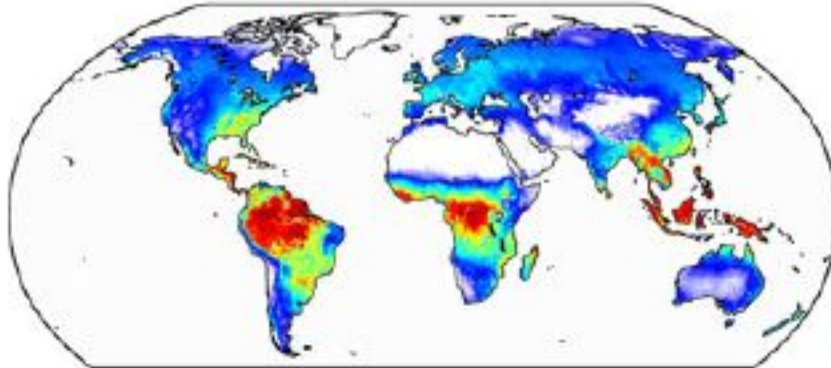
$$APAR = fAPAR * PAR$$

$$NPP = fAPAR * PAR * \epsilon$$

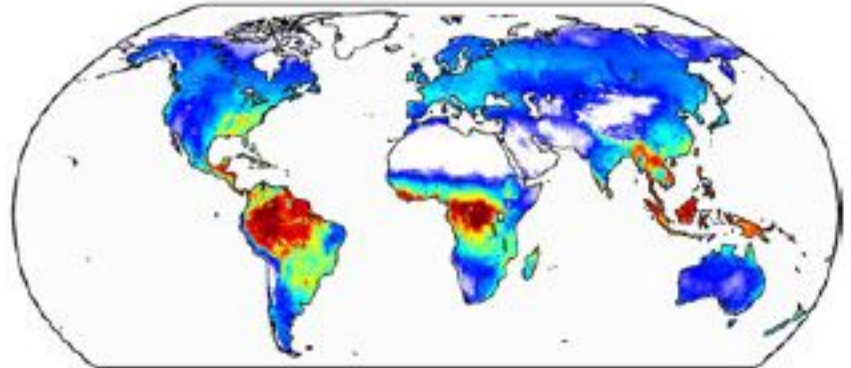
# global scale GPP estimate

GPP

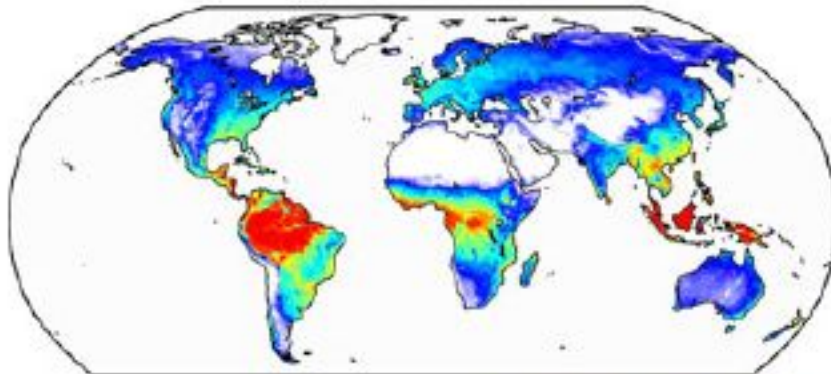
(a) ISAM-CRUNCEP



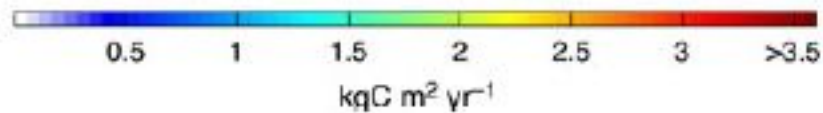
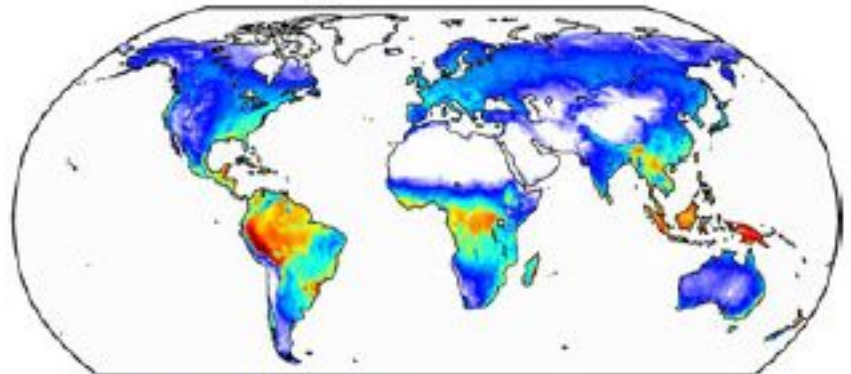
(b) ISAM-NCEP



(c) FLUXNET-MTE

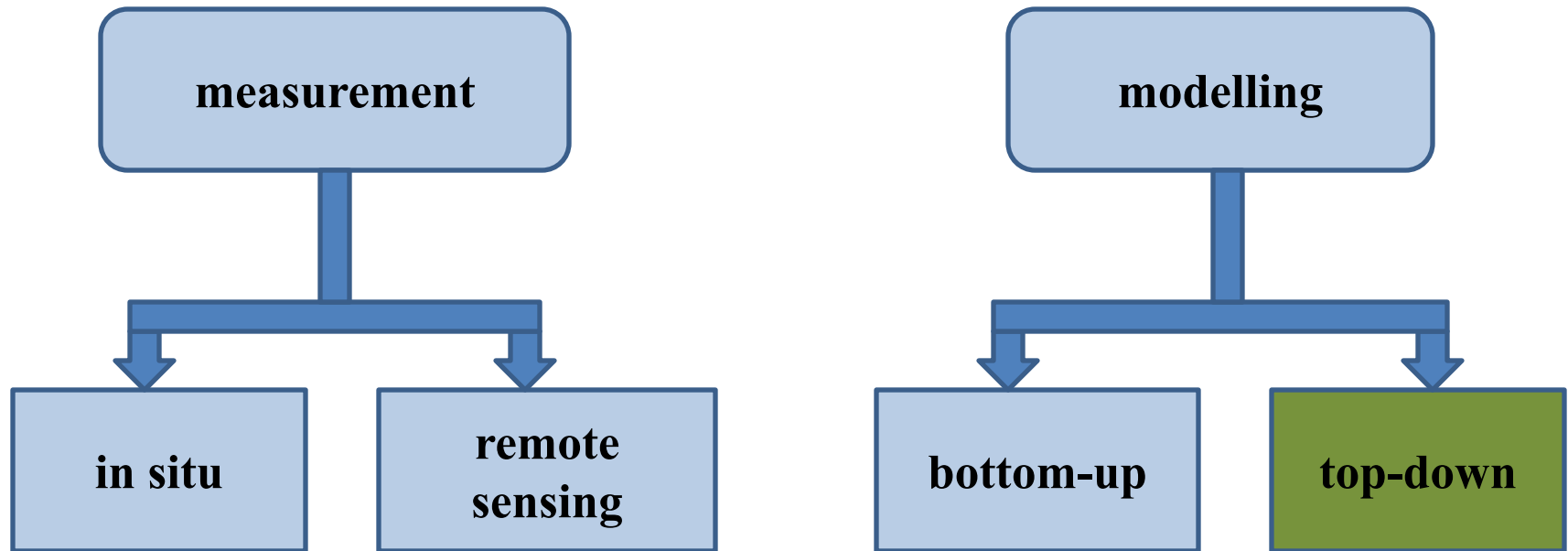


(d) MODIS-DAO



*Barman et al., (2014)*

# How to estimate the biosphere-atmosphere flux exchanges?



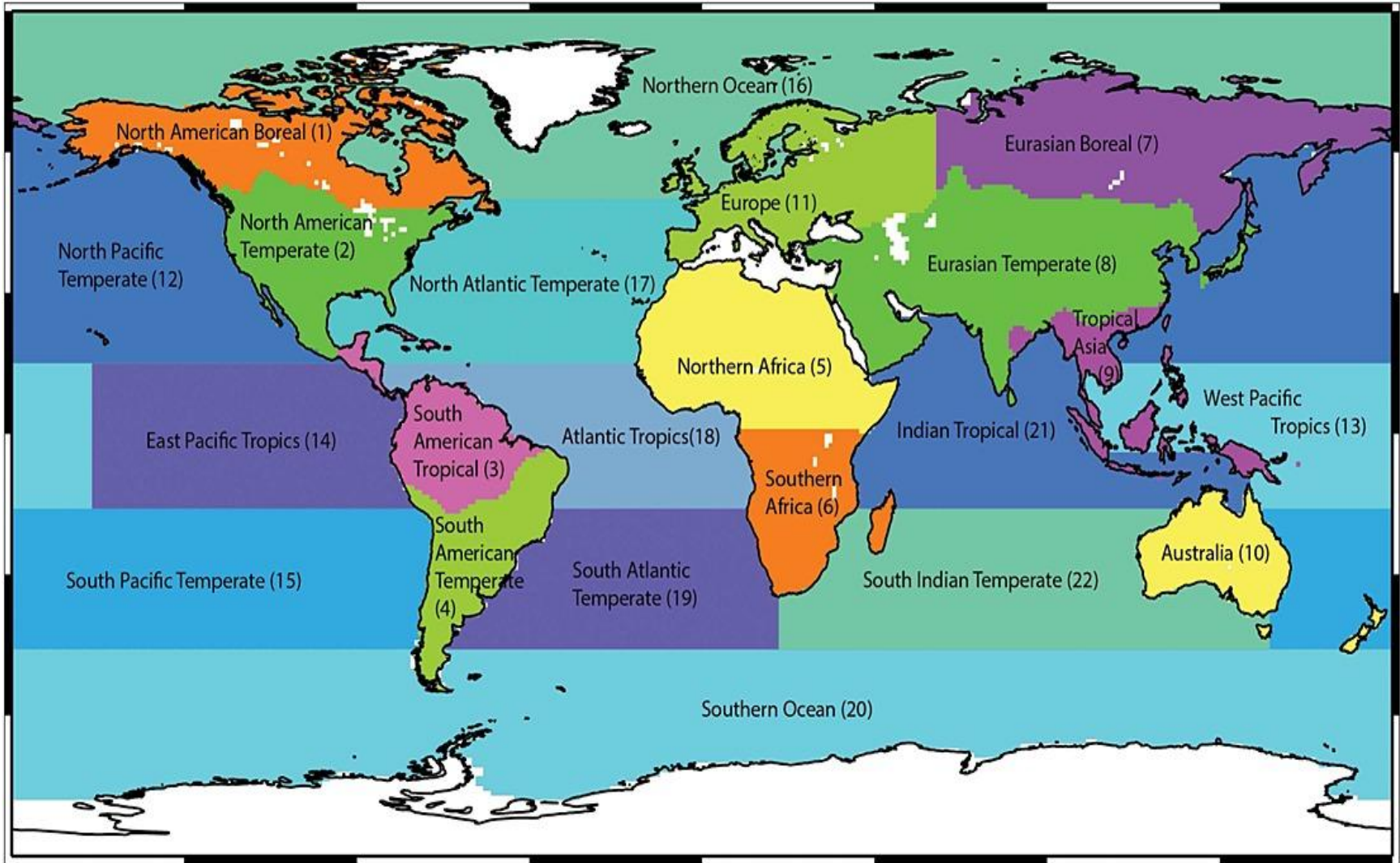
# Inverse Models

- the atmospheric CO<sub>2</sub> concentrations are observed at multiple spatial and temporal scales and the sources and sinks responsible for these distributions are computed by a backward modelling approach (Gurney et al., (2004); Rayner et al., (2005)).

- variables ( $\psi$ ) and observations ( $\chi$ ) matrices are connected by the following relation,

$$G \cdot \psi = \chi$$

- G is the coefficient matrix. The inverse of G is mostly calculated by Bayesian inversion (Heimann & Kaminski, (1999)).
- Observed spatio-temporal distribution of trace gases are partitioned into potential sources and sinks using atmospheric transport models (Kaminski et al., (1999)); Lagrangian (Pisso et al., (2019)), Eulerian (Pillai et al., (2012)) or hybrid schemes (Siqueira et al., (2000))



The different land (and ocean) zones used in the TRANSCOM inversion (Houweling et al., (2015)



# Robustness and uncertainty in terrestrial ecosystem carbon response to CMIP5 climate change projections

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nature  
geoscience

REVIEW ARTICLE

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## Terrestrial biogeochemical feedbacks in the climate system

A. Arneth<sup>1,2\*</sup>, S. P. Harrison<sup>3,4</sup>, S. Zaehle<sup>5</sup>, K. Tsigaridis<sup>6,7</sup>, S. Menon<sup>8</sup>, P. J. Bartlein<sup>9</sup>, J. Feichter<sup>10</sup>,  
A. Korhola<sup>11</sup>, M. Kulmala<sup>2</sup>, D. O'Donnell<sup>10</sup>, G. Schurgers<sup>1</sup>, S. Sorvari<sup>2</sup> and T. Vesala<sup>2</sup>

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# some interesting resources to explore

1. warning stripes :

<https://www.warningstripes.com/>

1. Eddy Covariance from NEON Science :

<https://www.youtube.com/watch?v=CR4Anc8Mkas>

2. UN REDD+ :

<https://www.un-redd.org/>

3. Intergovernmental Panel on Climate Change :

<https://www.ipcc.ch/>